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~5.9 cal ka BP Towada-Chuseri tephra from Towada volcano: A mid-Holocene marker layer from Japan to northeast China

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4 1 **~5.9 cal ka BP Towada-Chuseri tephra from Towada volcano: A**
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6 2 **mid-Holocene marker layer from Japan to northeast China**
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24 **Abstract**

25 Far-travelled ash layers from explosive volcanic eruptions can provide invaluable
26 marker horizons for dating and correlating regional to global sedimentary archives.
27 Here, we present a new cryptotephra associated with the ~5.9 cal ka BP Towada
28 Chuseri eruption (To-Cu) in a peat sediment record from northeast China. This tephra
29 exhibits a rhyolitic glass composition that can be distinguished from other widespread
30 tephra layers around the region of Japan and northeast China. Our findings extend the
31 known range of this ash significantly, making it now traceable about 1,200 km from
32 its source, Towada volcano, Japan. Notably, this tephra provides an important
33 isochron for synchronizing palaeoenvironmental studies during the mid-Holocene
34 period from the western Pacific, central Japan, Japan Sea, and northeast China.

36 **1. Introduction**

37 Volcanic ash can form regional to global time-equivalent marker horizons, and
38 thus provide a unique dating method for synchronizing studies of regional
39 palaeoclimatic and palaeoenvironmental archives (Lowe 2011). The development of
40 cryptotephra (invisible volcanic ash) analysis has extended the applications of
41 tephrochronology to ultra-distal regions several thousand kilometers away from the
42 volcanic vent, yielding trans-continental marker horizons (Bourne et al., 2016; Cook
43 et al., 2018; Jensen et al., 2014; Lane et al., 2013; Mackay et al., 2016;
44 Pyne-O'Donnell et al., 2016; Sun et al., 2014, 2021; van der Bilt et al., 2017). Japan
45 has one of the largest number of subaerial volcanoes of the Earth, and explosive
46 volcanic eruptions are very frequent here (Machida 2002). Tephras of Japanese origin
47 have been identified in ultra-distal regions; for example, Hokkaido and Tohoku
48 tephras have been reported from eastern North America and Greenland ice cores
49 (Bourne et al., 2016; Mackay et al., 2016). Sun et al. (2021) found tephras from the
50 ~7.3 cal ka BP Kikai and ~12.8 cal ka BP Sakurajima eruptions in Huguangyan Maar
51 Lake in southern China. However, whether Japanese tephras can be traced across
52 northeast China is still unknown, and due to limited applications of tephrochronology
53 in this region.

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4 54 In this study, we demonstrate the potential for long-distance sedimentary linkages
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6 55 through the identification of a Japanese tephra in Gushantun peat sediment from
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8 56 northeast China. Its glass shard composition indicates that corresponds to the ~5.9 cal
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10 57 ka BP Towada-Chuseri (To-Cu) eruption. The occurrence of distal To-Cu tephra in
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12 58 continental northeast China, >1,200 km away from the volcanic vent, implies that this
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14 59 tephra can be used to link palaeoclimatic records from central Japan's Lake Suigetsu
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16 60 to northeast China, and it can potentially be used as a mid-Holocene dating horizon
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18 61 within marine sediments from the Japan Sea.
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21 63 **2. Material and methods**

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23 65 2.1 Study site

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27 66 Our study site, Gushantun bog (42°18' N, 126°17' E; Fig.1), is located in the
28
29 67 Huinan Country of Jilin Province, western flank of the Changbai Mountains. The bog
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31 68 is about 1 km in diameter, and is surrounded by a basaltic platform. Peat sediments
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33 69 were collected from near the center of the bog using a Russian peat corer to a depth of
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35 70 6.45 m. The section of sediment used in this study (2.70– 4.20 m) consists of fibrous
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37 71 peat with some plant remains.
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39 72

40 73 2.2 Tephra separation

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42 74 The cryptotephra separation method developed by Turney (1998) and Blockley et
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44 75 al. (2005) was used in this study. First, contiguous 10 cm intervals were burned at
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46 76 550 °C for 4 hours, and then sieved through 80 and 30 µm meshes, retaining the
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48 77 fraction 30–80 µm. Secondly, 2.15 g/cm³ and 2.5 g/cm³ sodium polytungstate (SPT)
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50 78 was used to float glass shards from the sediments. A polarizing microscope was used
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52 79 to check for the presence of glass shards. Where shards were located, further samples
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54 80 were prepared at 1 cm sampling resolution; these samples were treated using the same
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56 81 approach to determine the peak concentration of glass shards. Further samples from
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58 82 levels corresponding to shard concentration peaks were treated using acid digestion
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60 83 (Dugmore et al., 1992) to remove organic materials in preparation for major element

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4 84 geochemical analysis. Samples were mounted on glass slides and covered in Buehler
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6 85 EpoxyCure2 resin, which were then ground and polished to expose fresh glass
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8 86 surfaces for electron microprobe analysis (EPMA).
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11 88 2.3 Geochronology

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13 89 Plant remains were hand-picked and submitted for accelerator mass spectrometry
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15 90 (AMS) ^{14}C dating to the Beta Analytic Radiocarbon Laboratory, Florida, USA (Table
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17 91 1; Fig. 2A, 2B). A Bayesian age-depth model was established using the P-sequence
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19 92 depositional model in OxCal 4.4 (Bronk Ramsey 2008; 2009) with the IntCal20
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21 93 calibration dataset (Reimer et al., 2020).
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24 95 2.4 Geochemical analysis

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27 96 We measured major and minor element oxide concentrations using wavelength
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29 97 dispersive spectrometer electron microprobe analysis (WDS-EPMA) on a JEOL JXA
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31 98 8100 at the State Key Laboratory of Lithospheric Evolution of the Institute of
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33 99 Geology and Geophysics, Chinese Academy of Sciences (IGGCAS), Beijing. Ten
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35 100 elements (Na, Mg, Al, Si, K, Ca, Fe, Ti, Mn, and P) were analyzed with an
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37 101 accelerating voltage of 15 kV, a beam current of 6 nA, and a beam diameter of 5 μm .
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39 102 The Na content was determined at the start of the analysis, and peak counting times
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41 103 used were 20 s for all elements except for Na (10 s). Secondary standard glass of
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43 104 MPI-DING including ATHO-G and StHs6/80-G were used to check the precision and
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45 105 accuracy of the data (Jochum et al., 2006). Data presented in figures have been
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47 106 normalized to an anhydrous basis.
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50 108 **3. Results and discussion**

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52 109 A cryptotephra layer with 160 shards/0.5g dry sediments was identified at a depth
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54 110 of 314 cm in the Gushantun core, and thus the tephra layer was labelled GST-314.
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56 111 This tephra layer is composed of colourless and vesicular glass shards (Fig. 2C), of
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58 112 which seven were analyzed. All glass shards exhibit a consistent rhyolitic
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60 113 composition, with SiO_2 concentrations ranging from 73.59 to 75.48 wt%, and low

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4 114 K₂O concentration (1.11 to 1.26 wt%) (Table 2, Fig. 3, Fig. 4). The age of GST-314 is
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6 115 constrained by the age model to between 6030 and 5676 cal a BP (95% probability).

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8 116 To identify the source of GST-314, we considered the geochemical compositions
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10 117 of tephra from the surrounding volcanic region. Gushantun bog is located in the
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12 118 Longgang volcanic field (LVF), and the composition of the volcanic products from
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14 119 this region is basaltic with no known silicic products (e.g., Fan et al., 2000; Liu et al.,
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16 120 2009; Sun et al., 2016). Therefore, the rhyolitic GST-314 tephra was not from a local
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18 121 eruption of the LVF. Changbaishan volcano on the border between China and North
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20 122 Korea, ~120 km to the west, experienced frequently felsic volcanic eruptions during
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22 123 the Quaternary period (Chen et al., 2016; Liu et al., 2015; Sun et al., 2014, 2017,
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24 124 2018; Yi et al., 2021; Zhang et al., 2018). However, Changbaishan tephtras
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26 125 predominantly with trachytic and rhyolitic glass compositions exhibit higher
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28 126 potassium content than GST-314, which excludes it as the source (Fig. 3). Ulleungdo
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30 127 is a Quaternary volcanic island in the Japan Sea that is known to have produced many
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32 128 explosive eruptions (Kim et al., 2014; McLean et al., 2018, 2020), including the U2
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34 129 tephra (5681–5619 cal a BP) which has a similar age to GST-314. With an even
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36 130 higher potassium content than the Changbaishan tephtras, Ulleungdo can also be ruled
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38 131 out as the source for the GST-314 tephra (Fig. 3).

39 132 Considering other coincident large explosive (volcanic explosivity index:
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41 133 VEI >5) eruptions from southeast to east Asia that have widespread tephra dispersals,
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43 134 there are also some potential source candidates for GST-314: the 7307–7196 cal a BP
44
45 135 Kikai-Akahoya eruption (K-Ah, Kikai volcano, Japan), the 6079–5004 cal a BP
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47 136 Pinatubo (Philippines) eruption, the 4894–4827 cal a BP SG14-0704 tephra recorded
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49 137 in Lake Suigetsu (unknown eruption), the 7340–7180 cal a BP KS₂ tephra from
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51 138 Ksudach caldera (Russia), the 6375–6220 cal a BP IAv12 (AV₄) tephra from
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53 139 Avachinsky volcano (Russia), the 7920–7620 cal a BP KHG tephra from Khangar
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55 140 volcano (Russia), and the 5986–5899 cal a BP Towada-Chuseri eruption (To-Cu,
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57 141 Towada volcano, Japan) (Ishimura and Hiramine 2020; McLean et al., 2018; Plunkett
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59 142 et al., 2015; Ponomareva et al., 2017; Portnyagin et al., 2020; Smith et al., 2013; Sun
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143 et al., 2021). Geochemically, the GST-314 tephra has a diagnostic glass composition

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4 144 with low potassium content and high silicon content which is consistent only with the
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6 145 To-Cu tephra and clearly separable from other broadly contemporary tephtras (Fig. 3
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8 146 and Fig. 4). Consequently, on the basis of comparable glass composition and eruptive
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10 147 age, we attribute GST-314 in Gushantun bog to the To-Cu eruption.

11
12 148 Towada is an active volcano located in northern Tohoku, Honshu Island,
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14 149 northeast Japan (Fig. 1). The Towada-Chuseri (To-Cu, ~5.9 cal ka BP) and Towada-a
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16 150 (To-a, AD 915) tephtras were produced by the largest volcanic events of Towada
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18 151 during the Holocene. The To-Cu eruption has been attributed a VEI of 5 based on a
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20 152 total estimated tephra volume of 9.18 km³ (Hayakawa, 1985; Newhall and Self 1982).
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22 153 To-Cu can be divided into three units: Chuseri pumice (Cu), Kanegasawa pumice
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24 154 (Kn), and Utarube ash (Ut), of which the Cu pumice is the lowermost bed and most
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26 155 widely distributed (Hayakawa, 1985; Ishimura and Hiramine 2020). While there is
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28 156 some overlap in their glass composition, Ishimura and Hiramine (2020) have shown
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30 157 that the units are separable on the basis of shard morphology and major element
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32 158 geochemistry, especially Ut, the youngest unit which is the product of phreatic
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34 159 eruptions. On the basis of previous investigations, To-Cu dispersal was mainly
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36 160 towards the south of Towada, and tephra has been identified in various sedimentary
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38 161 archives, the most distal occurrence being in Lake Suigetsu ~700 km southwest of the
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40 162 vent (Hirose et al., 2014; Ishimura et al., 2017; Kariya et al., 2016; McLean et al.,
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42 163 2018). The GST-314 tephra exhibits a homogeneous glass composition which cannot
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44 164 be separated from the Cu and Kn pumices and which is also similar to the records in
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46 165 Lake Suigetsu (Fig. 4), but it does not show the heterogeneous geochemistry of the Ut
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48 166 ash. Therefore, the GST-314 tephra is likely related to the pyroclastic fall deposits
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167 (Cu or Kn) of the To-Cu eruption rather than its late stage phreatic eruption (Ut).

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51 168 ¹⁴C dating has been widely used to constrain the eruptive timing of To-Cu: for
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53 169 example, charcoal within the tephra deposit yielded an age of 5390 ± 140 BP
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55 170 (GaK-9761, 6487–5898 cal a BP, 95% probability, calibrated using IntCal20)
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57 171 (Hayakawa 1983). Recently, the identification of the tephra in Lake Suigetsu has
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59 172 enabled a precise age of 5986–5899 cal a BP (95% probability) to be attributed to this
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173 eruption based on varve chronology combined with ¹⁴C dating (Bronk Ramsey et al.,

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4 174 2012; McLean et al., 2018; Staff et al., 2011). The precisely constrained Lake
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6 175 Suigetsu age makes this tephra a valuable age marker in sedimentary sequences.
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8 176 When we apply the Suigetsu age of To-Cu to the Gushantun age-model, age
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10 177 uncertainty is reduced considerably during this period (Fig 2B).

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12 178 During the Holocene, major tephtras reported from Japan Sea marine sediments
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14 179 are mainly visible beds (e.g., Furuta et al., 1986; Machida 2002), amongst which ash
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16 180 from ~7.3 cal ka BP K-Ah and AD 946 Changbaishan Millennium Eruption (B-Tm)
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18 181 are the most commonly reported layers (Furuta et al., 1986; Machida 2002). The
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20 182 presence of To-Cu in northeast China increases the known dispersal of this tephra to
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22 183 more than 1,200 km from the vent, and suggests that it too may have been deposited
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24 184 in the Japan Sea, but perhaps only as cryptotephra. The To-Cu tephra also was found
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26 185 in marine sediments from the Pacific coast of the Tohoku region (Ikehara et al.,
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28 186 2017). Thus, it can provide a high-precision mid-Holocene age marker linking the
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30 187 western Pacific Ocean, central and north Japan, very likely the Japan Sea, and
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32 188 northeast China (Fig. 1).

33
34 189 The isopachs of proximal To-Cu deposits show a mainly eastern distribution
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36 190 (Hayakawa 1985). However, identification of this tephra in distal sedimentary
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38 191 archives from central Honshu confirms a southwestern dispersal (Ikehara et al., 2017;
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40 192 Ishimura et al., 2017; Ishimura and Hiramine 2020; Kariya et al., 2016; McLean et al.,
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42 193 2018). The finding of To-Cu in northeast China now also shows that tephra was
43
44 194 transported to the west during this eruption. Most tephtras around northeast Asia have
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46 195 eastern dispersals, reflecting the prevailing westerly winds in the region (Lane et al.,
47
48 196 2017; Machida 2002). During the summer season, however, the Asian Summer
49
50 197 Monsoon brings wind from the Pacific Ocean towards the Asian continent. This
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52 198 weather pattern is consistent the south-westerly dispersal of To-Cu tephra.
53
54 199 Consequently, the eruptive season of To-Cu is unlikely to be winter, assuming a
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56 200 similar weather pattern as today.

57 201

58 202 **4. Conclusions**

59
60 203 Our discovery of the ~5.9 cal ka BP To-Cu tephra in Gushantun bog, northeast

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4 204 China, extends the known dispersal of this tephra to ~1,200 km west of Towada. Its
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6 205 presence in China suggests that it can be used as a mid-Holocene age marker horizon
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8 206 for various sedimentary archives in the wider region that will facilitate the dating and
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10 207 synchronization of mid-Holocene palaeoenvironmental records around northeast Asia.
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12 208 The tephra also raises the possibility that other tephra layers of Japanese origin may
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14 209 be detected here, emphasizing the potential for cryptotephrochronological
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16 210 applications across northeast Asia.

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4 355 **Figure 1.** Locations of the study site (Gushantun bog), Towada volcano, and major
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6 356 active volcanoes around northeast Asia. The black solid line represents the confirmed
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8 357 dispersal of Towada-Chuseri (To-Cu) tephra. The known distribution of other major
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10 358 Holocene tephras around this area are highlighted by dashed lines (data source: U-Oki
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12 359 and K-Ah is from McLean et al., 2018; QEA/B-Sg-08 is from Sun et al., 2018; B-Tm
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14 360 is generated from Chen et al., 2016, McLean et al., 2016, and Sun et al., 2015). The
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16 361 locations of other key sites mentioned in the text are also shown (SHL: Sihailongwan;
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18 362 YC: Yuanchi, SG: Suigetsu; CBS: Changbaishan).

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21 364 **Figure 2.** Tephrostratigraphy and age-depth model for the section of Gushantun peat
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23 365 discussed in this paper (A). Light grey shading is the number of glass shards per 0.5 g
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25 366 of dried sediment in 10 cm increments, and black bands refer to glass shard
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27 367 concentrations in the 1 cm samples. When the initial age for the GST-314 was found
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29 368 to be consistent with To-Cu, the age-depth model was re-run with the precisely
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31 369 constrained age for the To-Cu from Lake Suigetsu as determined by McLean et al.
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33 370 (2018) (B). Photos of the colourless glass shards from the GST-314 tephra (C).

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35 371
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37 372 **Figure 3.** Bivariable plots that compares the glass shard compositions of GST-314
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39 373 with the discussed known eruptions. Data sources: K-Ah, SG14-0704, Ulleungdo:
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41 374 McLean et al., 2018; Pinatubo: Sun et al., 2021; KHG: Cook et al., 2017; IAv12:
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43 375 Ponomareva et al., 2017; KS₂: Plunkett et al., 2015 and Ponomareva et al., 2017;
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45 376 Changbaishan: Sun et al., 2017, 2018.

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48 378 **Figure 4.** Biplots comparing glass shard composition of GST-314 recorded in
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50 379 Gushantun peat, with To-Cu recorded in Lake Suigetsu (SG14-0840; McLean et al.,
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52 380 2018), and three proximal units of To-Cu tephra from Japan (Ishiruma and Hiramine,
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54 381 2020).

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383 **Table 1.** Radiocarbon dates from plant remains from Gushantun peat. Radiocarbon
 384 ages were calibrated using the IntCal20 curve in OxCal 4.4 (Bronk Ramsey 2008,
 385 2009; Reimer et al., 2020).

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Depth (cm)	Lab code Beta-	$\delta^{13}\text{C}$ (‰)	^{14}C Age (yr BP)	Calibrated Age (cal a BP, 95% probability)
282	448787	-29.6	4490 ± 30	5297–5040, 5001–4986
311	448788	-27.7	5010 ± 30	5895–5805, 5794–5783, 5769– 5653, 5623–5607
408	448789	-27.3	7580 ± 40	8450–8327, 8234–8225

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Table 2. WDS-EPMA analyzed glass shards for the tephra from Gushantun peat (GST-314). All the data have been normalized to an anhydrous basis (original analytical totals shown). Comparative glass data of the To-Cu tephra in Lake Suigetsu is from McLean et al. (2018).

Sample	SiO ₂	TiO ₂	Al ₂ O ₃	FeO	MnO	MgO	CaO	Na ₂ O	K ₂ O	P ₂ O ₅	Analytical total
GST-314											
	74.67	0.43	13.34	2.47	0.09	0.61	2.68	4.44	1.18	0.08	99.94
	74.20	0.45	13.48	2.51	0.17	0.59	2.78	4.63	1.15	0.03	97.86
	74.73	0.40	13.33	2.46	0.20	0.59	2.80	4.18	1.24	0.08	98.34
	74.63	0.42	13.32	2.46	0.11	0.62	2.65	4.50	1.21	0.09	96.95
	74.18	0.48	13.45	2.42	0.10	0.65	2.73	4.57	1.26	0.17	100.63
	73.59	0.49	13.64	2.79	0.09	0.64	2.87	4.72	1.11	0.06	99.63
	75.48	0.44	13.21	1.88	0.07	0.50	2.54	4.59	1.23	0.05	98.72
ATHO-G											
Average (n=3)	75.39	0.24	12.29	3.33	0.10	0.09	1.68	4.23	2.64	0.01	
2σ	0.57	0.04	0.56	0.34	0.19	0.05	0.01	0.17	0.07	0.05	
Preferred value	75.60	0.26	12.20	3.27	0.11	0.10	1.70	3.75	2.64	0.03	
Uncertainty (95%)	0.70	0.02	0.20	0.10	0.01	0.01	0.03	0.31	0.09	0.00	
StHS6/80-G											
Average (n=3)	63.22	0.71	17.92	4.51	0.11	1.94	5.21	4.95	1.28	0.15	
2σ	0.28	0.06	0.40	0.14	0.04	0.05	0.08	0.51	0.04	0.10	
Preferred value	63.70	0.70	17.80	4.37	0.08	1.97	5.28	4.44	1.29	0.16	
Uncertainty (95%)	0.50	0.02	0.20	0.07	0.00	0.04	0.09	0.14	0.02	0.02	
Suigetsu Lake SG14-0840											
Average (n=25)	74.15	0.47	13.54	2.33	0.11	0.61	2.81	4.57	1.22	0.08	
2σ	0.45	0.06	0.36	0.23	0.11	0.12	0.21	0.30	0.10	0.05	

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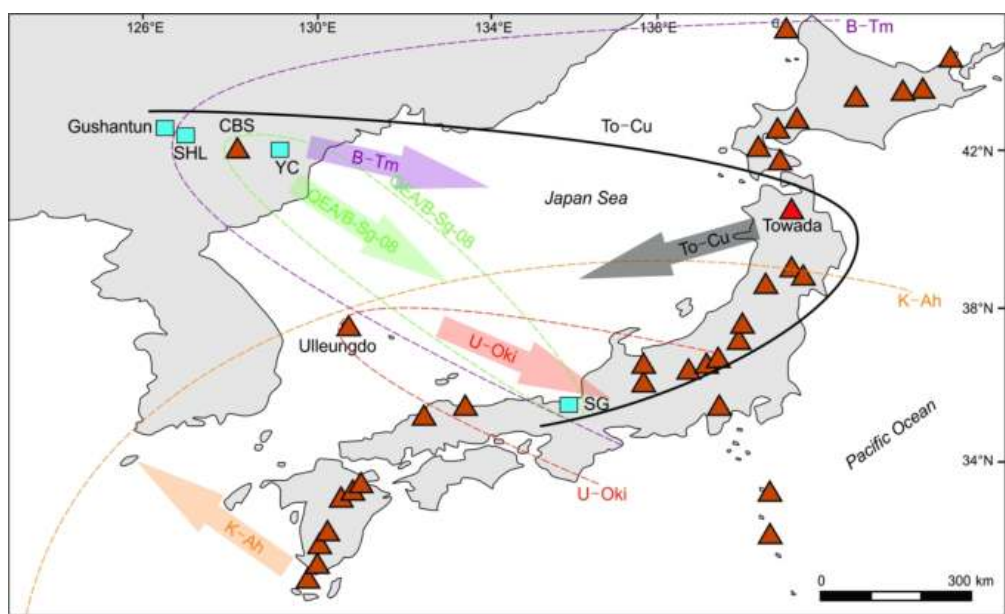


Figure 1

203x122mm (300 x 300 DPI)

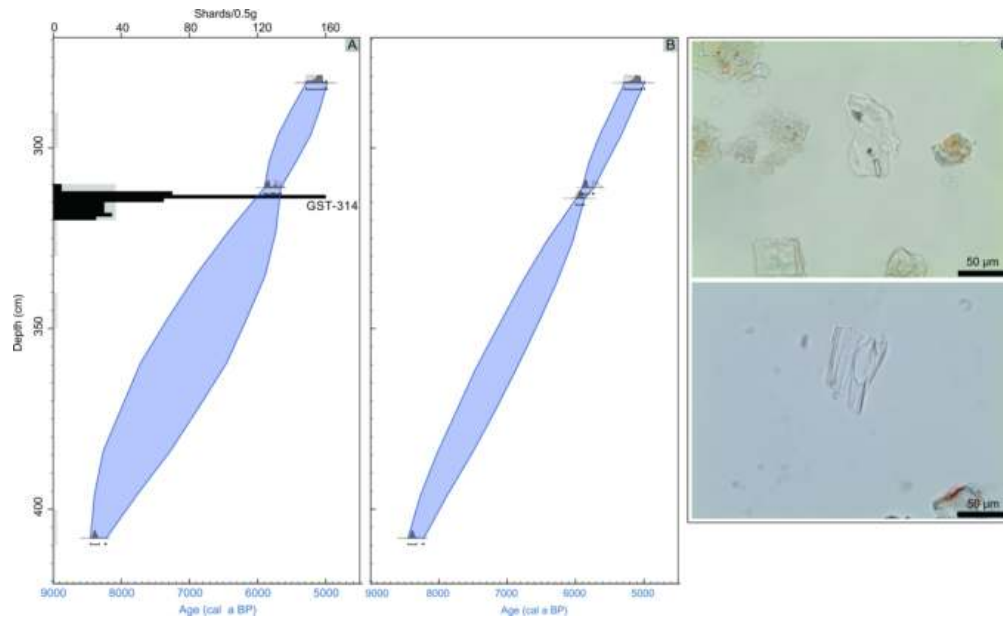


Figure 2

305x185mm (300 x 300 DPI)

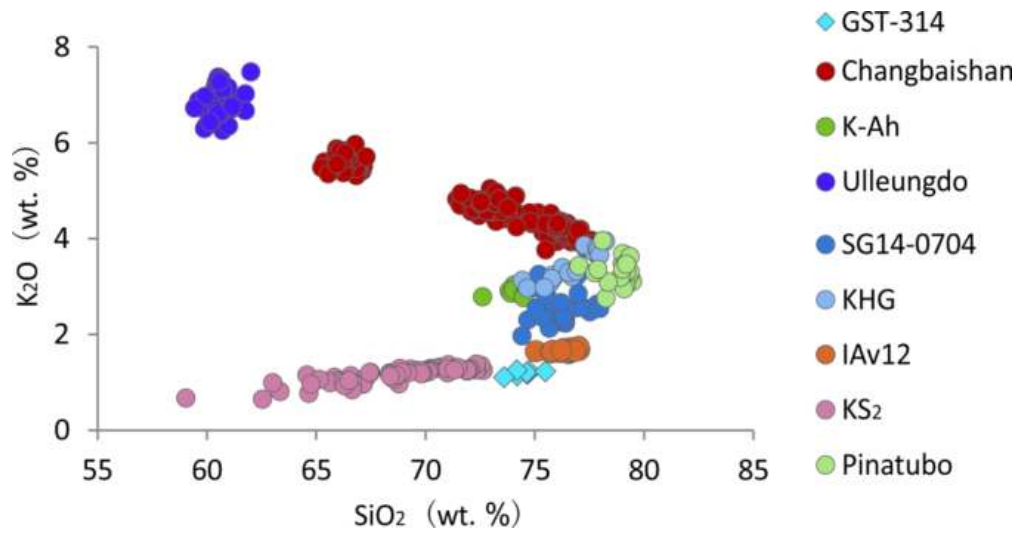


Figure 3

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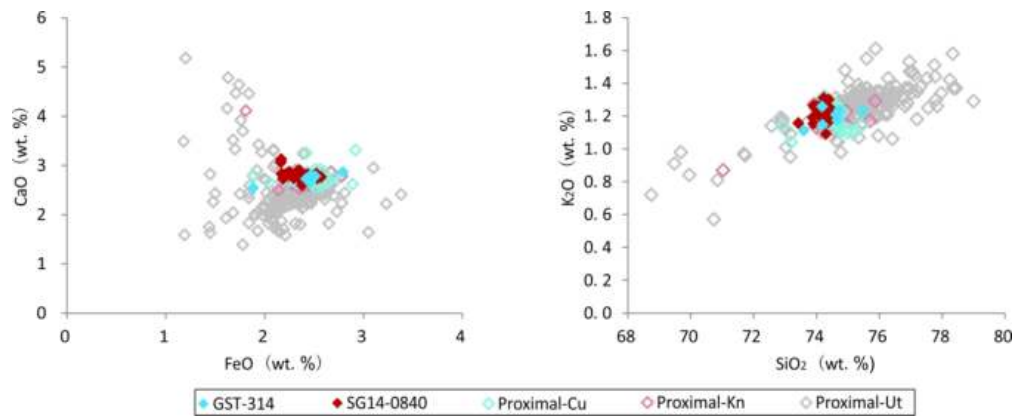


Figure 4

194x78mm (300 x 300 DPI)